Geostatistical-based geophysical model of electrical resistivity and chargeability data applied to image copper mineralization in the Ghalandar deposit, Iran

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ABSTRACT

This research aims to construct 3D geophysical models of electrical resistivity and induced polarization by interpolating 2D inverted physical models through the geostatistical approach. The applicability of the method was examined for the Ghalandar porphyry-skarn copper deposit in the Agh-Darah region, northwest of Iran. The 3D geophysical properties and block models of Cu grades were prepared by implementing the kriging interpolation method, whereby the recovered electrical models were closely linked to the Cu-sulfide mineralization. In order to evaluate the efficiency of the applied technique, the variogram models were validated using a cross-validation analysis of the kriging operation, proving the accuracy of data interpolation for each model. For the sake of meaningful correlation between geophysical models and Cu grades, the mineralization zones were extracted and subsequently propagated in the 3D space according to the generated physical properties. Meanwhile, the evaluation matrix was utilized to assess the performance of acquired results, where it confirmed that simultaneous consideration of physical models could much better determine the location of the copper mineralization. Also, the Swath plot was used as a second validation way to compare the anomalous zones.

Keywords: Chargeability, Copper Mineralization, Cross-Validation, Electrical Resistivity, Kriging

1. Introduction

Sulfide ore deposits are the main source of copper in the world, to which various geophysical surveys are applied to provide valuable information on the exploration of such sulfide-related targets [1]. Among the geophysical tools used in targeting these sources, the electrical properties of constituent materials with sharp contrasts are the most popular ones [2]. Constructing a 3D model of the blind target is the most important goal in geophysical explorations, while the model recovered from geophysical data inversion gives insights into the approximate geometry of the target. Electrical Resistivity Tomography (ERT), through a time-domain direct current survey, generates valuable information in close association with rock types and alterations [3]. Electrical resistivity (Rs) and induced polarization (IP) properties are usually derived from the ERT survey, which can delineate the location of sulfide-related copper mineralization. The integration of IP and Rs models has been successfully applied to detecting such sources [4]. The exploration of sulfides via Rs and IP methods is highly efficient due to electrical contrasts, and hence, the deposits with sulfides and oxides are generally identified by low resistivity and high chargeability values [5-10]. Laboratory measurements have shown that most of the sulfides represent IP effects that are larger than silicates and iron oxides [2]. However, IP responses are influenced by complex physical conditions [11-12].

Minerals and rocks associated with hydrothermal alterations in sulfide-bearing ore mineralization systems are recognized by distinct anomalous electrical properties. Also, geophysical methods, which are capable of localizing and modeling such electrical properties, are mainstays in the mineral exploration targeting of sulfide-related copper deposits. Indeed, electrical properties reflect the type and intensity of hydrothermal alteration related to these sources. The hydrothermal products that generate significant geophysical signatures are including pyrite, chalcopyrite, chalcocite, biotite, and sericite [13-14].

The dispersed nature of copper mineralization in sulfide-related systems is particularly suitable for IP surveys [15], initially developed for the exploration of porphyry copper deposits [16]. As a complex phenomenon, the IP anomalies reflect the ability of a rock to act as an electrical capacitor, where such characteristic is manifested in sulfide-bearing targets. Generally, the strongest IP anomalies are in association with quartz-sericite-pyrite alterations, mostly occurring in the sulfide-bearing systems [13-14]. The potassic alteration zone in the core is depleted of sulfide contents. The surrounding sericitic/phyllic alteration zone, however, has a higher sulfide content (e.g., pyrite), which itself is surrounded by the distal propylitic alteration zone with lower amounts of pyrite. In other words, the sericitic/phyllic alteration zone possesses the strongest IP anomalies in such mineralization systems [14].

Geostatistics, as a robust methodology in geospatial data interpolation, can add restrictions on spatial correlation, data conditioning, and incorporation of different scales [17]. The integration of geostatistics with geophysics has been widely used to tackling problems that arise from geophysical data modeling [4, 17-21]. Ramazi and Jalali (2014) investigated the application of geophysical inversion...
and geostatistical simulation to constructing electrical models on the basis of interpolating 3D physical properties from the 2D inverted ERT data. They used the IP and Rs models for qualitative and quantitative evaluations of a copper deposit, on which the Sequential Gaussian Simulation (SGS) was utilized to map the spatial distribution of physical models. Their results helped them optimize the location of exploratory boreholes [21]. Asghari et al. (2016) studied multivariate geostatistics based on a model of geo-electrical properties for copper grade estimation. They aimed to reduce the variance of estimation and related uncertainty. This method can be useful when a sporadic pattern of boreholes exists. In this study, the sulfide factor was incorporated as a secondary correlated variable to estimate Cu grade distribution. The results showed that the use of a secondary variable causes better results in comparison with ordinary kriging [22].

This work is focused on the Ghalandar sulfide-bearing copper deposit in the Eastern Azerbaijan province. The ultimate motive of research was to localize and extract Cu-bearing zones through a geophysical survey of the ERT, where those promising zones would be suggested for drilling to acquire information about the geometry of blind targets. The geostatistical tools were employed to the 2D inverted ERT data in order to construct the 3D models of electrical resistivity and induced polarization. Those recovered models were in close association with the 

Fig. 1 Structural geologic map of Iran, on which the location of the study area is presented in the NW portion, over the UDMA zone (modified after Richards et al. 2006) [30].

2. Geological Setting of the Ghalandar Porphyry-Skarn Deposit

The Northward Neo-Tethys subduction started in the Mesozoic, which led to the generation of the Iranian plateau [23-24]. Igneous activities in this subduction zone created a thick belt of mostly Cenozoic volcanic and plutonic units that is known as the Urmia-Dokhtar Magmatic Arc (UDMA), which is shown in Fig. 1. This belt has generated a distinct, linear intrusive-extrusive complex, located between Iran and parallel to the Sanandaj-Sirjan Metamorphic Zone (SSZ) and the central Iran microcontinent [25]. The Agh-Daragh mineral prospect, which is the case study in this work, is located within the UDMA zone (Fig. 2). The UDMA (also known as the Sahand-Bazman or Tabriz-Bazman zone) is the main host of many Iranian porphyry and epithermal Cu, Au, and Mo mineralization deposits [26-27]. This zone is about 50 to 100 km thick and mostly consists of an Andean-type magmatic arc adjoining the Central Iranian Micro-Continent (CIMC). In the structural geologic division map of Iran, the UDMA is characterized by the Cenozoic magmatic rocks of the Eocene-Quaternary age and their associated volcanoclastic rocks. Magmatic intrusions are mostly dominated by subvolcanic porphyritic granodiorite units of granite, granodiorite, diorite, and tonalite [28-29]. Fig. 1 illustrates the structural geologic map of Iran, on which the UDMA elongates from the NW to SE of the country. The Agh-Daragh area is located in the northwestern portion of the belt.

The simplified geological setting of the area is presented in Figs. 2 and 3a. In the western portions of the area and in the north of the granodiorite masses, there is a sequence of green tuffs and volcanic ashes related to the Cretaceous period. In the NE of the Gavdel village, there is a sedimentary unit, whose general color is dark gray and is generally composed of shale and limestone units. The main plutonic suite in the region, which is related to the Shiverdagh plutonic mass, is a porphyry granodiorite unit with pink feldspar phenocrysts [32].

Fig. 2 Field photo of different lithological units in the Agh-Daragh prospect zone, on which the location of the Ghalandar Cu mineralization (comprising the Ayran Goli and Gowdal mineralization systems) has been portrayed (Asgharzadeh Asl et al. 2017) [31].

3. Geophysical Electrical Survey

Considering the nature of Cu mineralization in the Ghalandar region, the ERT survey was carried out to delineate sulfide-related minerals. Scintrex IPR 12 equipment was used for data acquisition. Seven 2D ERT profiles with S-N directions were designed to acquire the information about the electrical resistivity and induced polarization properties of subsurface materials. An electrode spacing of 30 m with a line spacing of 100 m was conducted in the survey. The pole-dipole array was implemented based on the ERT configuration to inject the DC current into deeper locations. The layout of the ERT survey is indicated in Fig. 3b, showing that sixteen boreholes were drilled based on geophysical data modeling used to construct the geometry of sulfide-bearing Cu mineralization. The overall length of the drilling was about 2760 m.
4 presents a 3D visualization of all drilling with respect to the topographic surface, along with Cu grades.

(a) 3D visualization of all 2D inverted profiles of electrical resistivity (a), and chargeability (b).

Fig. 4. Location of boreholes with respect to the topographic surface in the study area. Cu grades are shown along each borehole as well.

The 2D inversion of the ERT data was carried out by utilizing the Res2dinv software developed by Loke, and the inverted sections for the electrical resistivity and induced polarization are visualized in Fig. 5. The size of geo-electric pseudo-section meshes was produced on the basis of the distribution of the data points as a rough guide. The height of the bottom row of the meshes is set to be equal to the equivalent height of investigating the data points with the maximum electrode spacing.

Inversion methods aim to determine the value of blocks that will produce an apparent resistivity pseudo-section that agrees with true measurements. The height of the first layer of the meshes for the pole-dipole array was set at 0.6 times that of the electrode spacing. The height of each subsequent layer was gradually increased by 10% [35]. The length of the blocks was equal to the electrode spacing in this study. The inverted models were such that the predicted data had the lowest misfit with original observations. Profile 4 was chosen as the representative of all ERT profiles to show the close consistency of drilled borehole neighboring the inverted section shown in Fig 6. The electrical models in profile 4 were estimated to the nearest borehole through the nearest neighborhood method to be able to compare them with Cu grades. Subsequently, the cross-correlation plots between Cu-Rs and Cu-IP were prepared, as displayed in Fig. 7. Pearson’s linear correlation coefficients between Cu and the electric variables were obtained to be equal to -0.76 and +0.81 for Rs and IP models, respectively. It should be mentioned that the higher values of Cu grades corresponded to lower and higher values of resistivity and induced polarization, respectively.

Fig. 5. 3D visualization of all 2D inverted profiles of electrical resistivity (a), and chargeability (b).

Since Cu enrichment was correlated with anomalous electrical models, constructing the 3D models of physical properties could provide insights into the geometry of the Cu mineralization in the Ghalandar region. Therefore, the following section discusses how geostatistical-based approach can facilitate the 3D recovering of electrical models from interpolating 2D inverted results shown in Fig. 5.
Fig. 6. A cross-section view of the 2D inverted models of the electrical resistivity (a), and the induced polarization (b) along profile 4, on which Cu grades along the drilled borehole next to the ERT survey are overlain on the sections.

Fig. 7. Scatter plots of Cu grades along the borehole near the ERT survey versus the electrical resistivity (a) and chargeability (b) for profile 4.

4. Geostatistical Models

This section discusses the procedure of constructing 3D models of two electrical properties along with the block model of Cu grades derived from exploratory boreholes. All boreholes in this study were drilled vertically. According to the distance between the boreholes, the block model dimensions were 20×20×10 m to generate the three aforementioned models. The statistical descriptions of the variables (i.e. Rs, IP, and Cu grade) derived from their composited models are summarized in Table 1. The histogram and box-plot of these variables are shown in Fig. 8, indicating a skewed distribution of each variable.

Since a spatial structure is important for implementing any geostatistical method, a variogram model is required to be determined in multiple directions. This tool can provide the main parameters for kriging estimations. Therefore, the accuracy of calculated variogram parameters has a substantial effect on the interpolated variable [36]. The directional variogram models for three variables were searched, and the highest spatial continuity for electrical resistivity, induced polarization, and Cu grade are shown in Fig. 9. Table 2 lists the main parameters of each videography, assuming a spherical model fitted to each variable.

Table 1. A summary of statistical properties of the variables: Rs, IP, and Cu grade.

<table>
<thead>
<tr>
<th></th>
<th>Number</th>
<th>Mean</th>
<th>Variance</th>
<th>Maximum</th>
</tr>
</thead>
<tbody>
<tr>
<td>Rs</td>
<td>606</td>
<td>755.98</td>
<td>2.17E+06</td>
<td>16676.1</td>
</tr>
<tr>
<td>IP</td>
<td>606</td>
<td>37.52</td>
<td>1310.46</td>
<td>172.85</td>
</tr>
<tr>
<td>Cu Sulfide</td>
<td>378</td>
<td>0.75</td>
<td>0.39</td>
<td>3.5</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th></th>
<th>Upper quartile</th>
<th>Median</th>
<th>Lower quartile</th>
<th>Minimum</th>
</tr>
</thead>
<tbody>
<tr>
<td>Rs</td>
<td>606</td>
<td>351.14</td>
<td>153.88</td>
<td>4.05</td>
</tr>
<tr>
<td>IP</td>
<td>49.12</td>
<td>26.66</td>
<td>12.23</td>
<td>0.04</td>
</tr>
<tr>
<td>Cu Sulfide</td>
<td>1</td>
<td>0.5</td>
<td>0.5</td>
<td>0</td>
</tr>
</tbody>
</table>

Fig. 8. The histogram and box-plot for the electrical resistivity (1st row), chargeability (2nd row), and Cu Sulfide content (3rd row).

In order to evaluate the accuracy of searched variogram models, a cross-validation method was taken into account to re-estimate the left-out data. Indeed, this validation method examines that the left-out data, and therefore, the desired statistical parameters must be reproduced by the estimated model. The comparison between the estimated and actual values shows the accuracy of the interpolation technique [37]. Here, a 2D cross-section of electrical properties and one borehole were left to be re-estimated by the aforementioned variogram model using the kriging method. The scatter plots of estimated versus original ones were presented in Fig. 10. Pearson’s linear correlation coefficients between actual and estimated values were obtained equal to 0.86, 0.92, and 0.77 for electrical resistivity, chargeability, and Cu grade, respectively. These values state that the variogram parameters provide sufficient accuracy for 3D modeling.
Fig. 9. Experimental directional semi-variogram models, and the number of pairs for (a) the electrical resistivity, (b) induced polarization, and (c) Cu grade.

Table 2. Variogram model parameters obtained for the three variables.

<table>
<thead>
<tr>
<th></th>
<th>Azimuth</th>
<th>Dip</th>
<th>Range</th>
<th>Sill</th>
<th>Nugget</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>IP</td>
<td>0</td>
<td>0</td>
<td>225</td>
<td>1300</td>
<td>20</td>
<td>spherical</td>
</tr>
<tr>
<td>Rs</td>
<td>180</td>
<td>20</td>
<td>117</td>
<td>210000</td>
<td>148706</td>
<td>spherical</td>
</tr>
<tr>
<td>Cu Sulfide</td>
<td>60</td>
<td>60</td>
<td>90</td>
<td>0.3</td>
<td>0</td>
<td>spherical</td>
</tr>
</tbody>
</table>

Fig. 10. Scatter plots of actual and estimated values for electrical resistivity (a), chargeability (b), and Cu sulfide content (c).
After determining the required inputs for implementing kriging, the 3D models of Cu grade, IP, and Rs were interpolated as well. The anomalous zones for each variable in the 3D model are visualized in Fig. 11. Also, the threshold values are determined by a multi-fractal analysis to plot the models. Discussing the procedure of implementing a fractal-based method is beyond the aims of this study.

In order to optimize the results and reduce the error, the IP and Rs anomalous zones were extracted simultaneously (Fig. 12), where both electrical models approve the probable location of Cu mineralization. In this case, the overall accuracy of the evaluation matrix increased to 0.6. The results show that the integration of IP and Rs has a better result for imaging Cu mineralization in comparison with their individual models. The integrated evaluation matrix is shown in Table 5.

### Table 4. Evaluation matrix of Rs and Cu estimations to test the performance of the 3D models.

<table>
<thead>
<tr>
<th>Cu mineralization Zone</th>
<th>Anomalous Rs block</th>
<th>Non-anomalous Rs block</th>
</tr>
</thead>
<tbody>
<tr>
<td>Inside</td>
<td>3446</td>
<td>12749</td>
</tr>
<tr>
<td>Outside</td>
<td>2724</td>
<td>16450</td>
</tr>
</tbody>
</table>

Another comparison approach is the swath plot, which is a graphical representation of a variable distribution from a series of slices assumed in three main directions of the block model [38]. For this purpose, the block model was divided into several slices in three main directions (NS, EW, and down-hole directions) and the average of block estimates (Cu, Rs, and IP) falling inside each slice was calculated. Finally, the average values were plotted against the corresponding direction (Fig. 13). It should be mentioned that the borehole and geophysical variables were normalized between -1 and 1 in order to compare the results better.

All plots were derived from block models shown in Fig. 12, which means all swath plots are related to the anomalous zone. As a general trend, the values of IP and Cu have a reverse trend in comparison with the Rs value, and they all fluctuate along the X-axis. The trends along the Y-axis are almost the same as the Y-axis.

### Table 5. Evaluation matrix of integrated IP-Rs and Cu grade estimations to evaluate the performance of 3D models.

<table>
<thead>
<tr>
<th>Cu mineralization Zone</th>
<th>Anomalous Rs and IP block</th>
<th>Non-anomalous Rs and IP block</th>
</tr>
</thead>
<tbody>
<tr>
<td>Inside</td>
<td>870</td>
<td>9175</td>
</tr>
<tr>
<td>Outside</td>
<td>229</td>
<td>12670</td>
</tr>
</tbody>
</table>

### 5. Conclusion

The performance of 3D electrical models was discussed for the Agh-Daragh porphyry-skarn copper deposit located in the Eastern Azerbaijan province of Iran. A geostatistical-based approach was utilized to interpolate the 3D geophysical models of electrical resistivity and induced polarization derived from the 2D inversion. The results of the kriging technique conducted on physical properties were in close association with the Cu model constructed from 16 boreholes in the Ghalandar prospect zone. Such correlation was confirmed by calculating the evaluation matrix between geophysical models and the Cu mineralization. The interesting result was obtained when the integrated
geophysical model had a higher correlation with the mineralized zones. The swath plots were carried out as the second validation method showing a high correlation between the anomalous zones.

![Swath plots](image)

**Fig. 13.** Normalized Swath plots of Cu, IP, and Rs values in the direction E-W (a), N-S (b), and vertical (c).

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